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092 A RELATED WORKS

094 With the evolution of scientific paradigms (Caño-Lores et al., 2020), researches on natural science
095 has mitigated from empirical and theoretical approaches to computational and data-intensive sci-
096 entific paradigms driven by computer simulations and big data (Babovi & Bajat, 2023; Li et al.,
097 2023). In the field of computational fluid dynamics, numerical models including FLUENT (Galeev
098 et al., 2013), MODFLOW (Brunner et al., 2010; Harbaugh, 2005; Hariharan & Shankar, 2017), and
099 COMSOL Multiphysics (Multiphysics, 1998), have achieved accurate simulations of physical pro-
100 cesses like complex multiphase flow, solute transport, and interactions between surface and soil wa-
101 ter within groundwater systems. Important issues in those spatial discretization methods for solving
102 the partial differential equations (PDEs) of groundwater are the computationally intensive. More-
103 over, the implementation of these models is typically challenging for personnel at water resource
104 management agencies, as highly specialized knowledge and extensive experience in numerical mod-
105 eling are required. Continuous advancements in data-intensive scientific paradigms have empowered
106 deep neural networks to achieve significant breakthroughs in image recognition (Krizhevsky et al.,
107 2017), meteorological modeling (Wan et al., 2018), material informatics (Niu et al., 2020), hydrology
(Maxwell et al., 2021), and mathematics (Choudhary et al., 2020), due to the powerful feature

108 extraction (Yu et al., 2013; Hinton & Salakhutdinov, 2006) and complex function approximation
 109 abilities (Cybenko, 1989). Nonetheless, as "black box" models, neural networks exhibit a lack of
 110 transparency in their decision-making processes and depend significantly on extensive training data,
 111 limiting their use in groundwater research. Recently, by combining physical laws with neural net-
 112 works, the physics-informed neural network (PINNs) method has improved the interpretability of
 113 models and diminished reliance on extensive datasets, providing an efficient and accurate approach
 114 for groundwater system modeling, and thus promoting the scientific and reliability of groundwater
 115 prediction and management decisions.

116 The concept of PINNs was first proposed by Raissi et al. (2019). In their approach, the neural net-
 117 work was trained by constructing physical constraints, and optimized by using the residuals based
 118 on initial conditions, boundary conditions, and governing equations as loss functions. Neural net-
 119 work derivatives were computed through automatic differentiation (AD). This method effectively
 120 solves both continuous and discrete PDEs, marking a new era of integration of physical knowl-
 121 edge and deep learning. The introduction of PINNs has significantly expanded the scope of neural
 122 networks in scientific research and engineering applications, while markedly enhancing the predic-
 123 tion accuracy, generalization ability, and interpretability of models. Since then, researchers have
 124 continued to explore and innovate within the PINNs framework, applying it to various fields such
 125 as fluid mechanics (Haghighat et al., 2021), structural dynamics (Moradi et al., 2023), acoustics
 126 (Song et al., 2021), and solid mechanics (Wang et al., 2023). In hydrological research, PINNs of-
 127 fer a new perspective for more convenient and efficient prediction of groundwater flow with their
 128 unique physical guidance characteristics. Wang et al. (2020) proposed a theoretically guided neural
 129 network model (TgNN) based on the loss function of groundwater seepage differential equations,
 130 which improved the generalization performance of the groundwater seepage model with limited
 131 observational data. Additionally, the team demonstrated that the physics-informed model could ef-
 132 fectively use physical information to predict groundwater seepage responses beyond the training set
 133 through groundwater seepage transfer learning tasks. Cho & Kim (2022) used an LSTM network
 134 to regress the residual differences between a neural network model and the Weather Research and
 135 Forecasting model-Hydro (WRF-Hydro). They developed a hybrid model (WRF-Hydro-LSTM) to
 136 predict groundwater levels, and experimental results indicated that this hybrid model achieved higher
 137 prediction accuracy, less sensitivity to training datasets, and better generalization performance than
 138 the LSTM network or WRF-Hydro model alone. Wang et al. (2021a) proposed a neural network
 139 constrained by geostatistical information and a physics-guided autoencoder based on convolutional
 140 neural networks (CNNs) (Wang et al., 2021b), applying the physical guidance method to the pa-
 141 rameter inversion of groundwater hydraulic conductivity fields. The model's accuracy and practi-
 142 cality were further enhanced. The study by (Pashaei Kalajahi et al., 2022) demonstrated that, even
 143 with limited data, physics-guided neural networks could accurately estimate key parameters such
 144 as hydraulic conductivity and porosity of the medium, showcasing their strong ability to combine
 145 data-driven approaches with physical guidance. Daolun et al. (2021) established a new specialized
 146 neuron model incorporating pressure gradient information and proposed an algorithm called sign-
 147 post neural network (SNN), which significantly improves the accuracy of solving unsteady seepage
 148 partial differential equations. Although significant progress has been made in improving groundwa-
 149 ter seepage models using PINNs, current models still heavily depend on the adequacy and quality
 150 of observed data and remain sensitive to outliers. If the training datasets are insufficient or contain
 151 corrupted data, the simulation results may exhibit considerable bias. Moreover, after the model has
 152 been trained, its applicability is generally limited to specific hydrogeological parameter settings,
 153 making it difficult to generalize to broader or unforeseen hydrogeological scenarios, thus restricting
 154 the model's versatility and practicality.

153 B BASIC LAW OF GROUNDWATER SEEPAGE

155 Without considering changes in water density, the flow of groundwater in a three-dimensional
 156 aquifer within porous media can be expressed by the following partial differential equation:
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$$158 \frac{\partial}{\partial x} \left(K_{xx} \frac{\partial H}{\partial x} \right) + \frac{\partial}{\partial y} \left(K_{yy} \frac{\partial H}{\partial y} \right) + \frac{\partial}{\partial z} \left(K_{zz} \frac{\partial H}{\partial z} \right) + W = \mu_s M \frac{\partial H}{\partial t} \quad (1)$$

160 Where t is time (T); K_{xx} , K_{yy} , and K_{zz} are the hydraulic conductivities (LT^{-1}) along the x , y ,
 161 and z axes, respectively. H is the hydraulic head (L) at the corresponding space-time point; W is

the source/sink term of groundwater (such as precipitation, pumping, etc.); μ_s is the specific storage (L^{-1}) of the aquifer below the free surface. M is the thickness of the aquifer. Combined with fixed solution conditions (boundary conditions and initial conditions), it can reflect the water balance relationship per unit volume and per unit time under Darcy flow conditions.

C DEEP NEURAL NETWORK

The Deep Neural Network (DNN) consists of an input layer, several hidden layers, and an output layer, where each layer contains multiple neurons connected by a weight matrix (which may include bias terms). Consider a deep neural network with L layers, denoted as $Y = f_{net}(X, \theta)$, with n^l neurons in the L -th layer. The number of neurons in the input layer equals the dimension of the input feature vector X , and the number of neurons in the output layer equals the dimensionality of the output vector Y . The value of the i -th neuron in L -th layer, $z_{l,i}$, is computed as the product of the weight $W_{l,i}^T$ and neurons from the previous layer, plus the bias term $b_{l,i}$:

$$z_{l,i} = W_{l,i}^T a_{l-1} + b_{l,i} \quad (2)$$

The output of neurons is nonlinearly transformed by the activation function, which provides nonlinear capability to the network. After processing $z_{l,i}$ with the activation function σ , the output of the i -th neuron in the L -th layer, $a_{l,i}$, is obtained:

$$a_{l,i} = \sigma(z_{l,i}) \quad (3)$$

The final output of the network is produced by the output layer of the multi-layer neural network, and the network training process relies on a loss function that quantitatively assesses the difference between the network's predictions and the actual target values.

D PERFORMANCE METRICS

To evaluate the performance of the model, three main indicators were used for quantitative analysis: Mean Absolute Error (MAE), Root Mean Square Error ($RMSE$), and Determination Coefficient (R^2). Their definitions are as follows:

$$MAE = \frac{1}{n} \sum_{i=1}^n |y_i - \hat{y}_i| \quad (4)$$

$$RMSE = \sqrt{\frac{1}{n} \sum_{i=1}^n (y_i - \hat{y}_i)^2} \quad (5)$$

$$R^2 = 1 - \frac{\sum_{i=1}^n (y_i - \hat{y}_i)^2}{\sum_{i=1}^n (y_i - \bar{y})^2} \quad (6)$$

Where n represents the total number of samples, y_i represents the reference output from MODFLOW of the i -th sample, \hat{y}_i represents the predicted value of the i -th sample by the DNN model, and \bar{y} is the average of the reference values of all samples. Ideally, the closer MAE and $RMSE$ are to 0, the smaller the prediction error and the higher the accuracy of the model. The closer the R^2 value is to 1, the better prediction the DNN model achieves.

E DETAILS OF PI-RGSM-K MODEL

Given that the PI-RGSM-K model is a specific application of the PI-RGSM model under conditions of a heterogeneous hydraulic conductivity field, and since the structures of both models are similar, most of the experimental settings are identical to PI-RGSM as follows:

(1) As the hydraulic conductivity field has been integrated into the physics-informed loss function, the model input features are spatial-temporal coordinates (x, y, t) , canal water levels (H_a, H_b) , and source/sink term W . The model still uses a random method to generate input data for self-supervised training (shown in Table 1).

(2) A heterogeneous hydraulic conductivity field is introduced, with K defined as a function that varies linearly with x (i.e., $K = -0.01x + 0.8$). The hydraulic conductivity linearly changes from $0.8m/d$ at the left boundary to $0.4m/d$ at the right boundary.

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Table 1: Input features and value ranges of the PI-RGSM-K

<i>Input features</i>	$x(m)$	$y(m)$	$t(d)$	$W(m/d)$	$H_a(m)$	$H_b(m)$
<i>Value ranges</i>	[0,40]	[0,10]	[0,100]	[0,0.007]	[2,3.5]	[2,3.5]